

Estimation and comparison of reference evapotranspiration using different methods to determine olive trees irrigation schedule in different bioclimatic stages of Tunisia

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Abstract. The study of olive trees water requirements allows a better water management by using more accurate methods including maximum parameters of the continuum soil-plant-atmosphere. The Penman-Monteith equations is considered as the most rational approach and the most reliable for calculating evapotranspiration. Only this approach necessarily requires an important number of climate parameters. The use of other equations, less complicated and using less climate parameters may be a reliable and efficient alternative. This experimental study was carried out on two cultivars cv. "Meski" and cv. "Chemlali" conducted in the intensive system in different bioclimatic stages (Subhumid, Semi-Arid and Arid) in Tunisia. This work aims to estimate olive trees water needs using evapotranspiration calculation in three different bioclimatic stages. For that, we compared the Penman-Monteith formula with Blaney-Criddel, Hargreaves-Temperature, Hargreaves-Radiation and Priestley-Taylor formulas to estimate reference evapotranspiration (ET_0). Results show that ET_0 values calculated by Priestley-Taylor and Blaney-Criddel formulas were more or less similar to Penman-Monteith. The ET_0 values found by Hargreaves-Temperature and Hargreaves-Radiation were twice the values calculated by Penman-Monteith formula. We also found good correlations between the reference evapotranspiration calculated by the Penman-Monteith equation and that calculated by Priestley-Taylor and Blaney-Criddel equations in all bioclimatic stages (R^2 more than 0.85, $p < 1\%$). The ET_0 sensitivity analysis has shown that solar radiation and air temperature (energetic climatic parameters) have the dominant effect on the ET_0 at the level of the different climatic regions. Accordingly, in the case of lack of some climatic parameters and in sub-humid, semi-arid and arid conditions and for the different phenological stages of the olive tree, we can use

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Priestley-Taylor and/or Blaney-Criddle formulas to estimate water needs.

Keywords: Reference evapotranspiration; Olive; Water requirement; Bio-climatic stages.

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Introduction

Climate change is imposing an additional burden on water management in areas where water resources are already strained in the rationale for the provisions (Levina, 2006). These changes are mainly concentrated in the northern hemisphere, particularly in the Mediterranean areas, which are known for their pronounced climate variability (Hedger and Cacouris, 2008). The temperature in these areas will increase with an average of 0.8 °C to 1.3 °C by 2020. In contrary, we will observe a decrease in precipitation from -5 to -8% concerning all seasons without exception (Stocker et al., 2013). Tunisia, located at the southern shore of the Mediterranean and forming part of these vulnerable zones, also presents a progressive scarcity of water resources essentially in the field of agriculture.

The olive tree is one of the most adapted species to water scarcity and drought condition. That is why the olive sector is a strategic sector in the Tunisian economy. Indeed, Tunisia's olive-growing heritage is estimated at more than 82 million trees, which cover an area of 1,835,000 hectares, representing around 30% of the agricultural area (DGPA, 2015; Jackson et al., 2015). Today, even if the olive oil sector continues to be competitive and plays an important role in the country's economy, several weaknesses and threats persist (Karray, 2012). Unfortunately, this sector suffers from the instability of production from one year to other due to inter and annual irregularity of the rains. The high

variability of production from one year to the next, significantly, affects the regularity of export flows, which causes significant fluctuations in the national and international markets. To cope with this situation, the use of olive trees in intensive mode is an efficient solution.

In this context, it is necessary to optimize irrigation by a real estimation of the water needs. According to Pereira et al. (2007) the sustainability of water resources depends on the technologies helping to decide: "When and how much irrigate". The estimation of water needs was carried out using different methods integrating the maximum parameters of the soil-plant-atmosphere continuum. Some research relates to the use of the climate method, which boils down to estimate reference and cultural evapotranspiration. Each of these methods is based on a set of climatic parameters.

A large number of formulas were used, developed and improved to calculate the water consumed in the form of daily or monthly reference evapotranspiration (ET_0) (Blaney and Criddle, 1950; Priestley and Taylor, 1972; Hargreaves and Allen, 2003; Bois et al., 2005; Summer and Jacobs, 2005; Temesgen et al., 2005; Alkaeed et al., 2006; Papova et al., 2006; Lovelli et al., 2008; Martinez and Thepadia, 2010; Todorovic et al., 2013; Raziera and Pereira, 2013; Masmoudi-Charfi and Habaieb, 2014; Bchir 2015). The Penman-Monteith equation (ET_0 -PM), which according to FAO (1998) is considered to be the most rational and reliable approach, requires an important set of climatic parameters. The use of

other equations that are less complicated and useless climate parameters can be a reliable and efficient alternative.

In order to find the least complicated and most adequate evapotranspiration calculation method in different experimental sites of study in Tunisia, we compared the Penman-Monteith equation (ET_0 -PM) with that of Blaney-Criddle (ET_0 -BC), Hargreaves Temperature (ET_0 -HT), Hargreaves Radiation (ET_0 -HR) and Priestley-Taylor (ET_0 -PT) equations.

Material and methods

Presentation of the experimental sites

The study was carried out in three experimental sites spread over three different bioclimatic stages; subhumid (El Hawaria), semi-arid (Enfidha) and arid (Gafsa), characterized by quite divergent climatic parameters and rainfall gradient from 173 mm to 557 mm/year respectively for the arid and subhumid regions.. At the level of each experimental field, a meteorological station was installed.

First experimental site at the Subhumibe (El Hawaria)

The study plot is an olive grove planted by the Meski variety, conducted

intensively with a spacing of 7 m x 7 m. It is characterized by the following geographic coordinates latitude 36° 53' N; longitude 10° 48' W, and altitude 93 m. During the year of the study (2010), rainfall recorded at this plot was 557 mm/year (Table 1). Evapotranspiration recorded in the same year is of the order of 1,514 mm/year (Table 1).

Second experimental site at the semi-arid level (Enfidha)

The study is carried out on an intensive olive variety 'Meski' a square spacing of 7 m x 7 m. It has the following geographical features: altitude 23 m; longitude 10° 22' E; latitude 36° 08' N. At this site and during 2010; rainfall was 295.4 mm/year, and evapotranspiration was in the order of 1,482 mm/year (Table 1).

Third experimental site at the Aride (Gafsa)

The study plot is an olive grove planted by the variety Chemlali conducted in intensive with a spacing of 4 m x 4 m. The geographic coordinates of the site are: latitude 34° 28' N; longitude 5° 50' E and altitude 350 m. Rainfall recorded during 2010 and at the level of the study plot was 173.5 mm/year (Table 1). Evapotranspiration was in the order of 1,727 mm/year (Table 1).

Table 1. Environmental conditions (air temperature, relative air humidity, rainfall and evapotranspiration) during the study period at the three experimental fields (National Meteorological Institute, 2010).

	Air temperature (°C)			Relative air humidity (%)			Rainfall (mm/month)			Evapotranspiration (mm/month)		
	Sub-humid (El Hawaria)	Semi-arid (Enfidha)	Arid (Gafsa)	Sub-humid (El Hawaria)	Semi-arid (Enfidha)	Arid (Gafsa)	Sub-humid (El Hawaria)	Semi-arid (Enfidha)	Arid (Gafsa)	Sub-humid (El Hawaria)	Semi-arid (Enfidha)	Arid (Gafsa)
January	11.8	11.9	10.9	77	69	60	90.1	34.2	20.2	67	65	78
February	13.1	13.3	13.3	77	65	49	79.2	42.0	0.6	70	76	94
March	15.1	15.0	16.2	82	72	47	65.3	27.8	0.1	95	99	132
April	16.0	17.1	19.7	83	78	54	43.2	38.8	52.8	109	107	154
May	19.1	20.1	22.0	76	63	44	19.9	12.0	2.1	160	160	190
June	23.5	23.7	27.6	74	68	41	3.0	1.2	4.8	182	200	215
July	27.5	28.2	30.9	73	62	37	0.0	0.0	0.5	240	241	234
August	27.6	27.3	30.7	72	61	40	0.0	0.8	0.0	205	194	216
September	23.4	24.4	25.6	81	68	50	64.1	25.8	7.7	133	124	152
October	19.8	18.9	20.3	82	76	53	95.6	70.8	32.9	108	88	115
November	16.1	16.5	14.3	80	74	60	68.9	36.4	35.2	78	68	77
December	13.2	11.6	11.6	78	70	56	27.7	5.6	16.6	67	60	70
Total							557	295.4	173.5	1.514	1.482	1.727

Measurement of climatic parameters *in situ*

Evapotranspiration is conditioned by many factors such as solar radiation, air temperature, relative air humidity, atmospheric pressure and wind speed. The meteorological factors are easily measurable. Indeed, the global solar radiation and the heat flux in the soil are measured regularly every 30 min. For simultaneous determination of air temperature and humidity, a psychrometer is used. The meteorological station placed in the experimental plot also includes an anemometer for wind speed measurements, which is installed 2 m above the canopy, at a place

free from any obstacle (tree). The rain gauge is the basic instrument of precipitation measurement. It indicates the global rainfall precipitated in the interval of time separating two readings (once a day).

Evapotranspiration calculation methods

To calculate ET_0 , there are a lot of empirical formulas using different climatic parameters (Masmoudi-Charfi and Habaieb, 2014; Bchir, 2015). The use of the different empirical formulas depend on available weather data. Climatic data allow us to calculate evapotranspiration using the following formulas:

Method	Equation	Used parameters
Blaney-Criddle (1950)	$ET_0 = P \times (0,46 t + 8)$	t: Monthly mean air temperature (°C). P: the average daily percentage of annual diurnal hours as a function of latitude.
Hargreaves Température (Bois et al., 2005)	$ET_0 = 0,0023 \times R_A \times TD^{0.5} \times (T + 17,8)$	R_A : Extraterrestrial radiation ($cal/cm^2 \cdot day^1$); TD: difference between maximum temperature and minimum temperature (°C); T: Monthly mean air temperature (°C)
Hargreaves Radiation (Bois et al., 2005)	$ET_0 = 0,0135 \times (T + 17,8) \times R_s$	T: Daily mean air temperature (°C). R_s : solar radiation (MJ/m^2).
Priestley-Taylor (Hargreaves and Allen, 2003)	$ET_0 = \alpha \frac{\Delta}{(\Delta + \gamma)} (R_n - G)$	α : an empirical correction for our case we used $\alpha = 1,26$ (Sumner and Jacob, 2005), R_n : net radiation (W/m^2), G: soil heat flux (W/m^2), the slope of the vapour pressure curve ($kPa/°C$); γ : psychrometric constant ($kPa/°C$).
Penman-Monteith (Allen et al., 1998)	$ET_0 = \frac{0.408 \Delta (R_n - G) + \gamma \frac{900}{T + 273} U_2 (e_a - e_d)}{\Delta + \gamma (1 + 0.34 U_2)}$	R_n : net radiation at the surface of culture ($MJ / m^2 \cdot day$); G: soil heat flux ($MJ / m^2 \cdot day$); T: mean air temperature at 2 m (°C); U_2 : wind speed measured at 2 m (m/s); $(e_s - e_a)$: vapor pressure deficit (kPa); Δ : the slope of the vapour pressure curve ($kPa/°C$); γ : psychrometric constant ($kPa/°C$) and 0.34: coefficient of the wind (s/m);

Results and discussion

Calculation of reference evapotranspiration (ET_0)

Figure 2 reports the results of the application of different methods to estimate ET_0 in different bioclimatic stages (subhumid at El Hawaria; semi-arid at Enfidha and arid at Gafsa). The values obtained show that the reference evapotranspiration varies considerably according to the calculation method. While admitting that the Penman-Monteith formula is the reference method. The variation in the difference between the values recorded by the Penman-Monteith method and the other methods generally depends on the variation of the climatic factors involved in each formula.

The comparison between the different methods at the three study stations shows that the evapotranspiration values calculated by the methods of Hargreaves Temperature (ET_0 -HT) and Hargreaves Radiation (ET_0 -HR) are the most overestimating methods compared to values found by Penman-Monteith formula (ET_0 -PM). The work carried out in Portugal by Paredes and Rodrigues (2010), showed that there is a small error when using the HT equation with respect to that of PM, this error is more important at the level of wetlands compared to arid regions. According to Alexandris et al. (2008) and Martinez and Thepadia (2010), in a humid climate, the equation of HT overestimates the ET_0 with respect to the ET_0 -PM.

The Priestley-Taylor (ET_0 -PT) and the Blaney-Criddle (ET_0 -BC) methods give ET_0 values similar to those calculated by the Penman-Monteith formula (ET_0 -PM) throughout the study year and the 3 bioclimatic studied stages (Figure 2). Similar results have found by Alexandris et al. (2008). This can be explained by the fact that solar radiation, which constitutes the energetic part of both formulas, is the major factor controlling evapotranspiration (Xiaoying and Erada, 2005; Bchir, 2015). Generally, the variation of values recorded by the Penman-Monteith method and the other methods depends on the variation of the climatic factors involved in each formula. The PM method has more accurate values compared to other methods, because of its large number of parameters involved in the formula. On the other hand, the other formulas are limited to the global solar radiation and/or the average air temperature, which generates values of less precision. Also, the overestimation of the ET_0 recorded by the other methods (other than PM) could be explained by the use of the average data in the formula which considerably increases the variation of the ET_0 (Baldy, 1998), especially in the hot period of the year (High radiation and temperature). In the same context, Bouhlassa and Paré (2006) found that the Blaney-Criddle and Hargreaves methods overestimate the reference evapotranspiration compared to that calculated by the Penman-Monteith method.

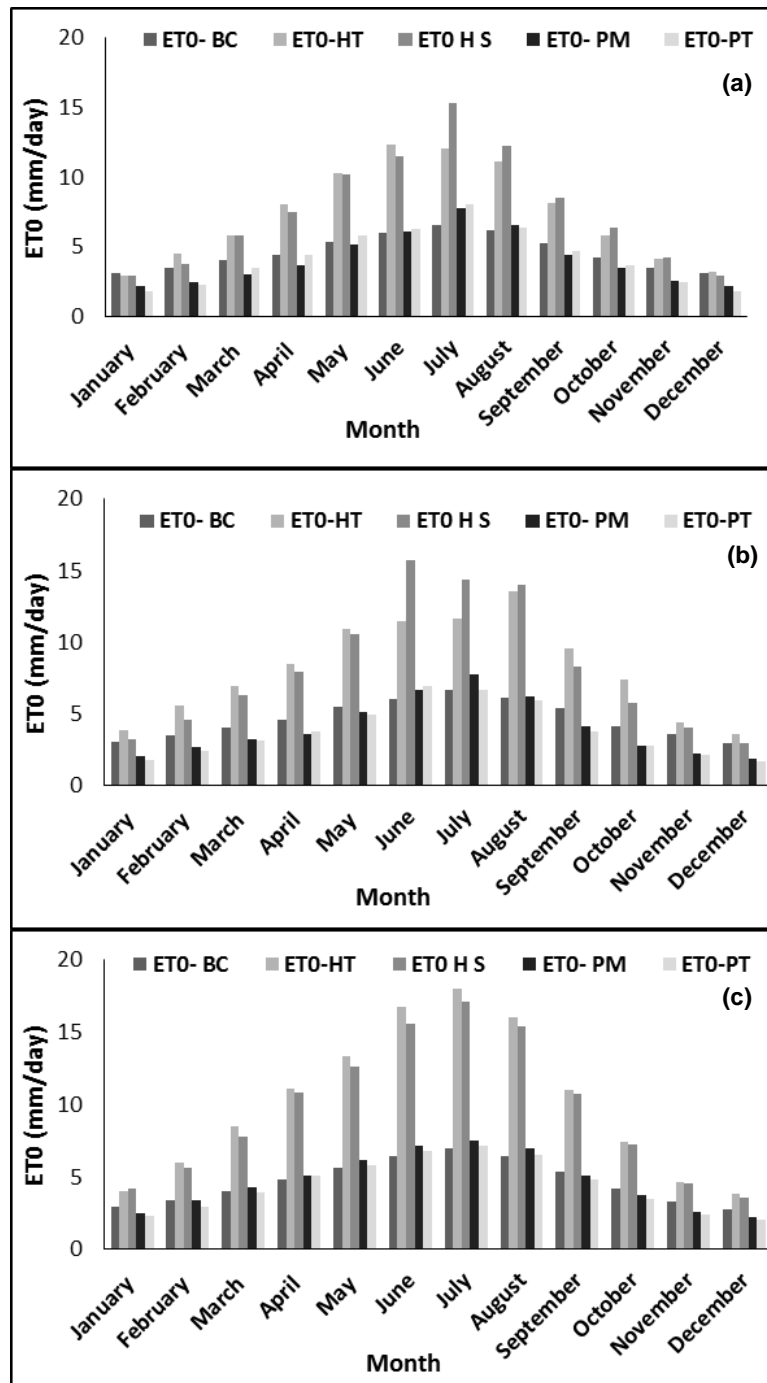


Figure 2. Reference evapotranspiration (ET_0 , mm/day) values estimated by several methods: Penman-Monteith (ET_0 -PM), Blaney-Criddle (ET_0 -BC), Hargreaves Temperature (ET_0 -HT), Hargreaves Radiation (ET_0 -HR) and Priestley-Taylor (ET_0 -PT) during the year 2010. (a) El Hawaria: Subhumid, (b) Enfidha: Semi-arid, and (c) Gafsa: Arid.

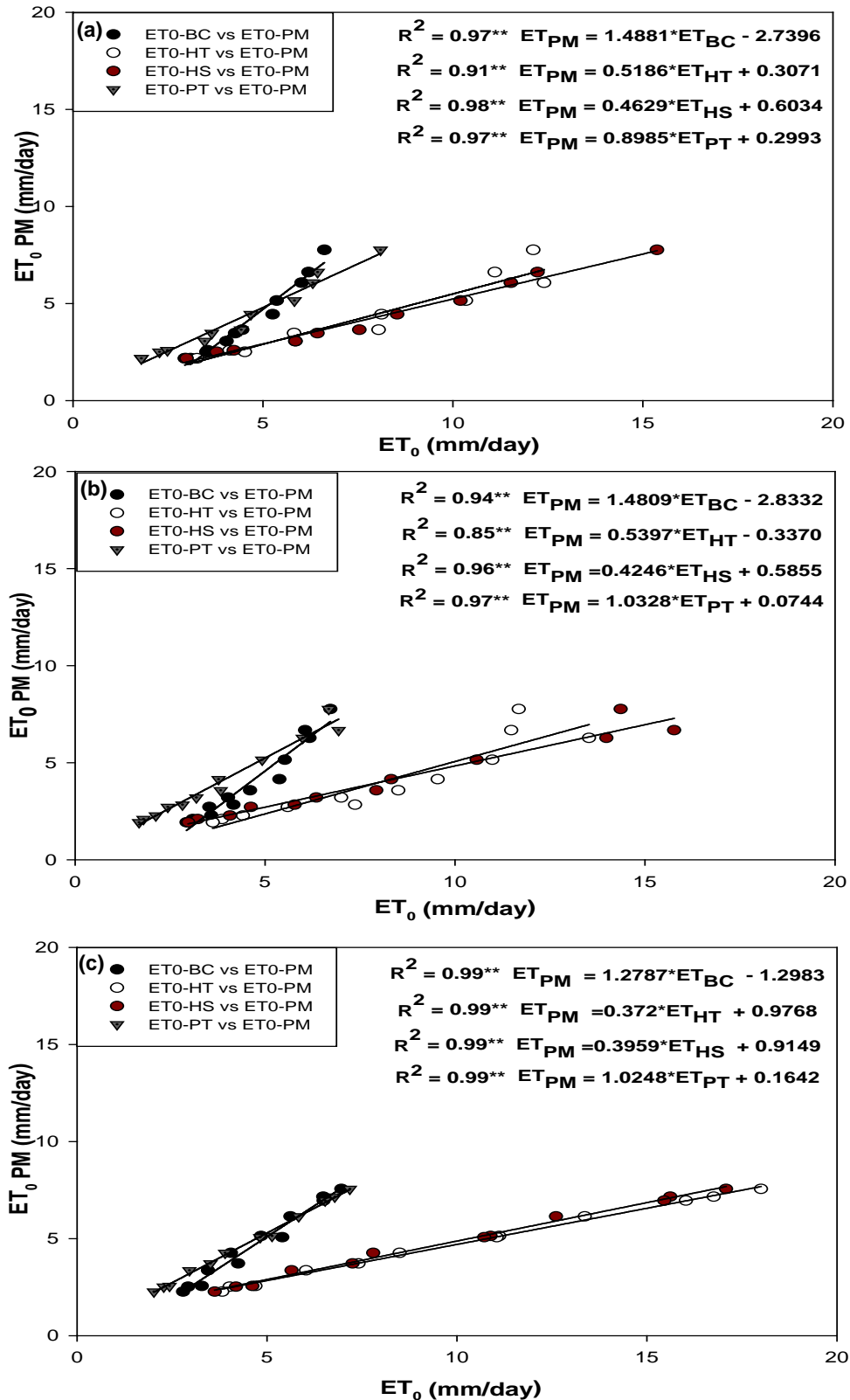


Figure 3. Linear correlation between of ET_0 calculated by the method of Penman-Monteith and several methods: Blaney-Criddle (ET_{0-BC}), Hargreaves Temperature (ET_{0-HT}), Hargreaves Radiation (ET_{0-HR}) and Priestley-Taylor (ET_{0-PT}). (a) El Hawaria: Subhumid, (b) Enfidha: Semi-arid, and (c) Gafsa: Arid.

Correlations between the different calculating formulas

In order to understand the difference between the used methods, the linear regression ($ET_0\text{-PM} = ax + b$) was studied between the ET_0 values calculated at the 3 bioclimatic stages (Subhumid, Semi-arid, Arid). The effectiveness of the empirical methods with respect to the Penman-Monteith formula ($ET_0\text{-PM}$) was evaluated in all the study sites in the graphical representation of Figure 3. Highly significant and positive correlations are observed between the Penman-Monteith and the various other formulas, with R^2 values exceeding 0.85 ($p < 1\%$). By analyzing the linear regression equations, the Priestley-Taylor formula gives the ET_0 values closest to those found by the Penman-Monteith formula. According to Priestley and Taylor (1972), it is sufficient to have solar radiation for the determination of evapotranspiration. Bois et al. (2007) showed that the ET_0 calculated by the Priestley-Taylor formula has a highly significant correlation with the ET_0 calculated by the Penman-Monteith method ($R^2 = 0.94$, $p < 1\%$). The Blaney-Criddle (BC) equation also gives similar ET_0 values (with a lowestimation) to those recorded by the Penman-Monteith (PM) equation. In the case of lack of climate data, Masmoudi-Charfi and Habaieb (2014) found that the Blaney-Criddle formula is positively correlated with the Penman-Monteith equation in different bioclimatic stages of Tunisia. However, the same authors concluded that, at the phenological scale of the olive tree, this formula is effective only during the vegetative rest, the vegetation development, the beginning of the flowering, the fruit growth, the oil synthesis and ripening of olives.

ET_0 sensitivity study to climatic parameters

Variation in climatic parameters directly and differently affects

evapotranspiration (ET_0) (Goyal, 2004; Gong et al., 2006; Gao et al., 2016). This variation has been studied at the three experimental stations (El Hawaria, Enfidha and Gafsa). A variation of 10% for the different climatic parameters, involved in the calculation of ET_0 by the Penman-Monteith formula, has been established. At the El Hawaria (Subhumid) station, the increase in the relative air humidity by around 10% in 2010 decreases the ET_0 in a clear way (Figure 4). Indeed, this decrease varies from 7.7% (July) to 18.4% (December). The effect of solar radiation increase is month dependent. The highest increases values in ET_0 followed by an increase of 10% in solar radiation are observed from April to November. The same evolution was found in air temperature but with less importance. The wind speed did not show any significant effect concerning other climatic parameters. ET_0 varies within a range of -0.7% to 1.5% due to an increase (10%) in wind speed.

At the Enfidha (semi-arid) station, an increase of the radiation (10%) induces an average increase of 7.5% in ET_0 (Figure 4). The effect of radiation is less important during the winter months (about 6.4%). From March to August, the effect of air temperature is considered as the second important parameter. The effect of relative air humidity decreases in summer but it is remarkable during the rest of the year, with an average varying from -6% to 10%. The wind speed, at the semi-arid level as well as at the subhumid level, does not show a significant effect in comparison with the other climatic factors on the ET_0 variations.

In the Arid Region of Gafsa, net radiation is the most important factor in the ET_0 calculation (between 6.4% and 7.6%) throughout the year (Figure 4.c). The air temperature shows a side effect on the increase of the ET_0 , concerning the radiation. Rising of the relative air humidity decreases the ET_0 value by

about -3.8% during the winter months. In comparison with other stations (subhumid and semi-arid), the relative air humidity affects ET_0 less during the

rest of the year and more specifically during the summer months. The effect of the wind speed (U_2) on the ET_0 does not exceed 1.6% for a 10% increase in U_2 .

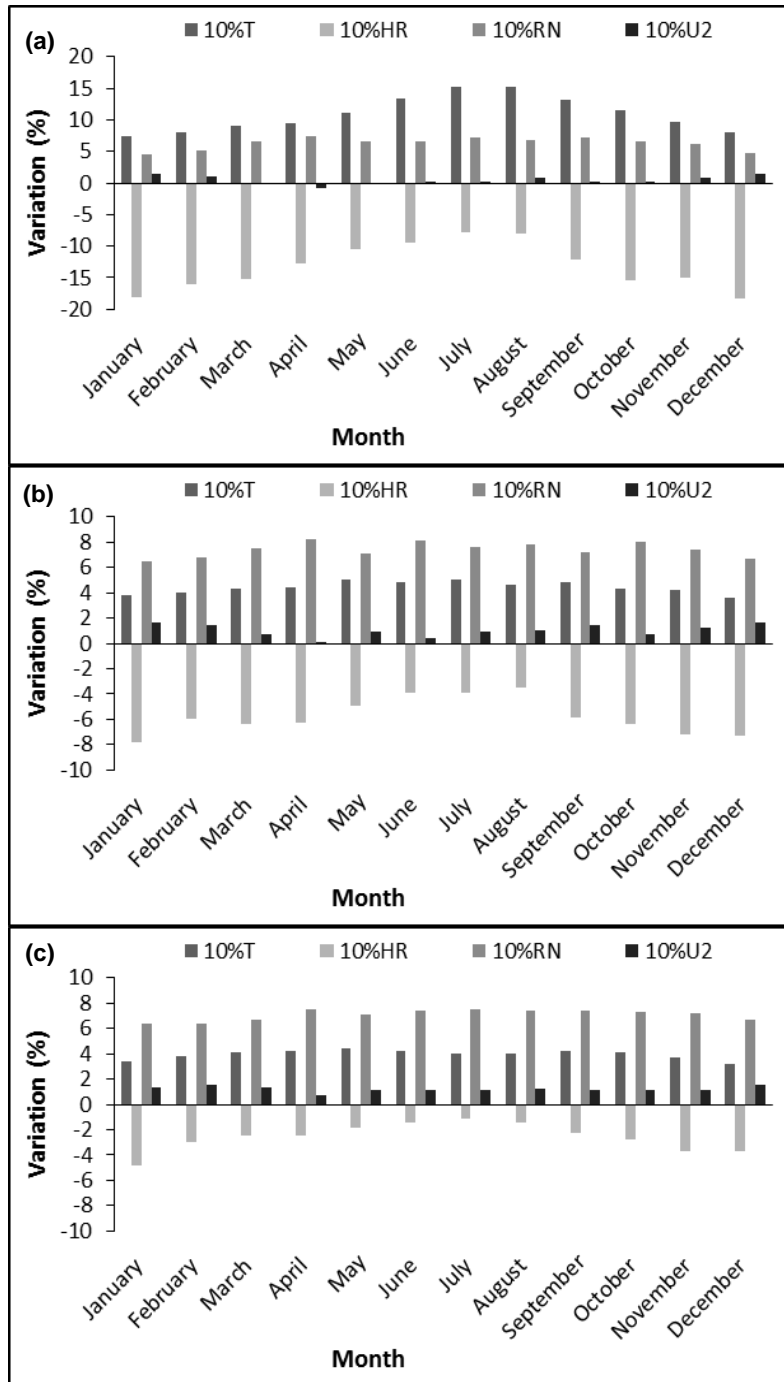


Figure 4. Reference evapotranspiration sensitivity to climatic parameters. (a) El Hawaria: Subhumid, (b) Enfidha: Semi-arid, and (c) Gafsa: Arid. T: Air temperature; HR: relative air humidity; RN: Net radiation and U_2 : Wind speed.

Analysis of the correlations between the reference evapotranspiration (ET₀) and the different climatic parameters

To understand the relationships between the different ET₀ calculation methods, it is important to see the effect of climate parameters and their interactions with ET₀. The correlation coefficients between the ET₀ and the different climatic parameters studied are shown in Table 2. Results showed that at different bioclimatic studied stages, the ET₀ is positively and significantly correlated with the net radiation (R² = 0.97, p < 1%), the mean air temperature (R² = 0.91, p < 1%) and the vapor deficit pressure (R² = 0.79, p < 1%). Whereas, the relative air humidity shows

a negative and significant correlation (R² = -0.35, p < 1%) with ET₀. According to various authors, the most affecting parameters on variations of evapotranspiration are air temperature, humidity, wind speed and sun radiation (Wrachien and Mambretti, 2015; Yannopoulos et al., 2015; Khoshravesh et al., 2017; Valipour, 2017).

These results show also that the dominant parameter of ET₀ is solar radiation, which constitutes the energetic part of the Penman-Monteith equation. Several authors have similarly found the high sensitivity of the Penman-Monteith formula to changes in solar radiation (Bois et al., 2007; Xu et al., 2006; Yang et al., 2009; Tabari and Hosseinzadeh-Talae, 2014).

Table 2. Matrix of correlations between ET₀ and the different climatic factors.

	Tm	HR	RN	U ₂	VPD	ET ₀
Tm	1,00	-0,31*	0,83**	0,02 ^{NS}	0,78**	0,91**
HR		1,00	-0,33*	0,69**	-0,79**	-0,45**
RN			1,00	0,14 ^{NS}	0,65**	0,97**
U ₂				1,00	-0,39**	0,04 ^{NS}
VPD					1,00	0,79**
ET ₀						1,00

** Highly significant correlation (p < 1%); * Significant correlation (p < 5%) and NS: Not significant correlation. Tm: Mean air temperature; HR: relative air humidity; Rn: Net radiation; U₂: Wind speed and VPD: Vapor pressure deficit).

Olive tree water requirements at different phenological stages

The comparison between the different calculation formulas of the ET₀ and the analysis of the interaction between the ET₀ and the climatic parameters give us the possibility to select the most efficient formulas to estimate of the water requirements for olive tree. The determination of water requirements was based on the FAO formula (ET_c = Kr * K_c * ET₀, Fereres and Castel, 1981).

In our study, the cultural coefficient (K_c) varies between 0.46 and

0.65 (Paster et al., 1998; Bchir, 2010) and the reduction coefficient (Kr) retained for the experimental period was about 0.75 (Bchir, 2010; Elsayed-Farag, 2014). The appropriate formula for calculating ET₀ varies according to the bioclimatic stage and the phenological stage of the olive tree (Table 3). The Priestley-Taylor formula is the most efficient compared to Penman-Monteith one during the majority of phenological stages, at the three experimental sites (subhumid, semi-arid and arid). For some stages, it is also possible to use the Blaney-Criddle formula. At the level of the subhumid and

at the time of flowering (May), this formula gives results close to those found by the Penman-Monteith formula. It also gives E_{Tc} values comparable to those estimated by the Penman-Monteith formula in the first place and by that of Priestley-Taylor second, for the flower bud stages (mid-March and April), fruiting and fruit development (June). The Priestley-Taylor formula also keeps its performance in the semi-arid (Enfidha), with the possibility of using the Blaney-Criddle formula from the flowering stage to the fruit growth stage. In the region of Gafsa (Arid), the use of

the Blaney-Criddle formula differs on the time scale compared to the first two stations. From the flower induction (February) to the flowering stage (May), it is possible to use this formula.

Masmoudi-Charfi and Habaieb (2014) found that the Blaney-Criddle formula gives estimates of water requirements in the regions of Nabeul and Sousse. In the region of Sidi Bouzid, characterized by a climate similar to that of Gafsa, Masmoudi-Charfi and Habaieb (2014) showed that the formula of Turk, which is based on radiation and temperature was the most adequate.

Table 3. Table 3. Calendar of water requirements of the olive tree based on the most appropriate methods for the estimation of ET₀ and determined by phenological stages at the different bioclimatic stages.

	Phenological stage	Vegetative rest	Flower induction	Resumption of flowering	Appearance of flower buds	Flowering	Fructification	Fruit development	Core hardening	Fruit growth			Beginning of maturation		Complete maturation		Total (mm/year)
		January	February	March	April	May	June	July	August	September	October	November	December				
Subhumid El Hawaria	E _{Tc} calculating method	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor /Blaney-Criddle	Blaney-Criddle	Blaney-Criddle/ Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor		
	E _{Tc} -PM (mm/month)	25	26	34	39	55	75	96	91	53	38	38	25	595			
	Adopted E _{Tc} (mm/month)	21	24	39	47 / 47	57	74 / 78	100	88	56	40	36	21	603			
Semi-arid Enfidha	E _{Tc} calculating method	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor/ Blaney-Criddle	Priestley-Taylor/Blaney-Criddle	Blaney-Criddle/ Priestley-Taylor	Blaney-Criddle/ Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor		
	E _{Tc} -PM (mm/month)	24	29	36	38	55	83	96	86	49	31	33	22	582			
	Adopted E _{Tc} (mm/month)	21	26	36	41	53 / 59	86 / 75	83 / 82	85 / 82	45	31	31	20	555			
Arid Gafsa	E _{Tc} calculating method	Priestley-Taylor	Blaney-Criddle	Blaney-Criddle /Priestley-Taylor	Priestley-Taylor/ Blaney-Criddle	Priestley-Taylor /Blaney-Criddle	Priestley-Taylor/Blaney-Criddle	Priestley-Taylor /Blaney-Criddle	Priestley-Taylor /Blaney-Criddle	Priestley-Taylor /Blaney-Criddle	Priestley-Taylor /Blaney-Criddle	Priestley-Taylor	Priestley-Taylor	Priestley-Taylor			
	E _{Tc} -PM (mm/month)	29	35	48	54	66	89	93	95	60	41	37	26	674			
	Adopted E _{Tc} (mm/month)	27	36	45 / 44	54 / 51	62 / 60	84 / 80	88 / 86	90 / 89	58 / 64	38	36	24	642			

Conclusions

Adequate and correct estimation of ET₀ is so important in agricultural and hydrological studies, water resources and watershed management. In particular, it is necessary to support irrigation scheduling. Our results showed that the Priestley-Taylor equation (ET₀-PT) seems to reduce the risk of ET₀ overestimation, followed by the Blaney-Criddle equation (ET₀-BC). The formulas of Hargreaves Temperature (ET₀-HT)

and Hargreaves Radiation (ET₀-HR) overestimate ET₀. The sensitivity of analysis of ET₀ to different climatic parameters shows that the climatic energetic parameters (R_n and T_m) have the dominant effect on the ET₀. This explains the strong correlations found between the ET₀ calculated by the Penman-Monteith formulas (ET₀-PM), Priestley-Taylor (ET₀-PT) and Blaney-Criddle (ET₀-BC). This allowed the determination of olive tree water requirements scheduling based on the

most appropriate methods and corresponding to each phenological stages at different bioclimatic regions.

As a result, in the case of lack of certain climatic parameters and in Tunisian sub-humid, semi-arid and arid conditions, the use of the Priestley-Taylor equation and/or the Blaney-Criddle equation for estimating water requirements is easier than the rest of equations. This could be important for more economical management of water inputs, particularly in the irrigated public and private areas.

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Conflict of interests

Authors declare that there are no conflicts of interest.

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